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Helicopter-borne magnetic, electro-magnetic and radiometric geophysical survey in Okstindan and Røsvatnet area, Nordland county.



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Summary:

NGU conducted an airborne geophysical survey in Hemnes and Hattfjelldal municipalities, as part of NGU's general airborne mapping program. The data acquisition in the Okstindan area was started and completed in August 2021.

This report describes and documents the acquisition, processing and visualization of the acquired datasets and presents them in maps. The geophysical surveys consist of 2039 line-km data, covering an area of 394.5 km² flown from the base at Vesterli Skistadion in Korgen, Hemnes

The NGU modified Geotech Ltd. Hummingbird frequency domain EM system supplemented by an optically pumped Cesium magnetometer and the Radiation Solutions 1024 channels RSX-5 spectrometer mounted on a AS350-B3 helicopter was used for data acquisition.

The survey was flown with 200 meters line spacing, azimuth 90°, average speed was 90 km/h, average height clearance of the bird was 65 m and 93 m for the spectrometer.

Collected data were processed at NGU using Geosoft Oasis Montaj software. Raw total magnetic field data were corrected for diurnal variation and leveled using Geosoft micro-levelling algorithm. Radiometric data were processed using standard procedures as recommended by International Atomic Energy Association (IAEA).

EM data were filtered and leveled using both automated and manual levelling procedures. Apparent resistivity was calculated from in-phase and quadrature data for three coplanar frequencies (880Hz, 6.6kHz and 34kHz), and for two coaxial frequencies (980Hz and 7kHz) separately using a homogeneous half space model.

All data were gridded using cell size of 50x50 meters and presented as 40% transparent grids with shaded relief on top of topographic maps.

Keywords:	Airborne	Geophysics
Magnetic	Gamma spectrometry	Radiometric
Electromagnetic	Technical report	

CONTENTS

1. SURVEY SPECIFICATIONS.....	5
1.1 Airborne survey parameters.....	5
1.2 Airborne survey instrumentation	6
1.3 Airborne Survey Logistics Summary	7
2. DATA PROCESSING AND PRESENTATION.....	8
2.1 Total Field Magnetic Data	8
2.2 Electromagnetic data.....	10
2.3 Radiometric data.....	11
3. PRODUCTS	16
4. REFERENCES.....	16
Appendix A1: Flow chart of magnetic processing	17
Appendix A2: Flow chart of EM processing	17
Appendix A3: Flow chart of radiometry processing.....	17

FIGURES

Figure 1: Helicopter survey area in Okstindan and Røsvatnet area.....	4
Figure 2: Pegasus helicopter with pilot and NGU Hummingbird EM system	7
Figure 3: Gamma-ray spectrum with K, Th, U and Total Count windows.....	11
Figure 4: Okstindan and Røsvatnet survey area with flight path. The red segments were flown above 150m.	19
Figure 5: Total Magnetic Field	20
Figure 6: Magnetic Horizontal Gradient	21
Figure 7: Magnetic Vertical Gradient	22
Figure 8: Magnetic Tilt Derivative	23
Figure 9: Apparent resistivity. Frequency 7000 Hz, Coaxial coils	24
Figure 10: Apparent resistivity. Frequency 6600 Hz, Coplanar coils	25
Figure 11: Apparent resistivity. Frequency 980 Hz, Coaxial coils	26
Figure 12: Apparent resistivity. Frequency 880 Hz, Coplanar coils	27
Figure 13: Apparent resistivity. Frequency 34133 Hz, Coplanar coils.....	28
Figure 14: Radiometric Total Counts	29
Figure 15: Potassium ground concentration	30
Figure 16: Uranium ground concentration	31
Figure 17: Thorium ground concentration	32
Figure 18: Radiometric Ternary Image	33

TABLES

Table 1. Flight specifications	4
Table 2. Instrument Specifications	6
Table 3. Hummingbird EM system, frequency, and coil configurations	7
Table 4. Survey specifications summary.....	7
Table 5. Specified channel windows for the 1024 RSX-5 system.	11
Table 6. Maps in scale 1:100000, available from NGU on request.	16

INTRODUCTION

In 2021 NGU received government funds to acquire airborne geophysical data from parts of Nordland county. The survey area covers two areas with Okstindan mountain-area in the north, and further south an area between Røsvatnet and the Swedish border. The helicopter survey data presented in this report, amounts to 2039 line-km, or 394.5 km², with the area covered shown inside the red lines in Figure 1.

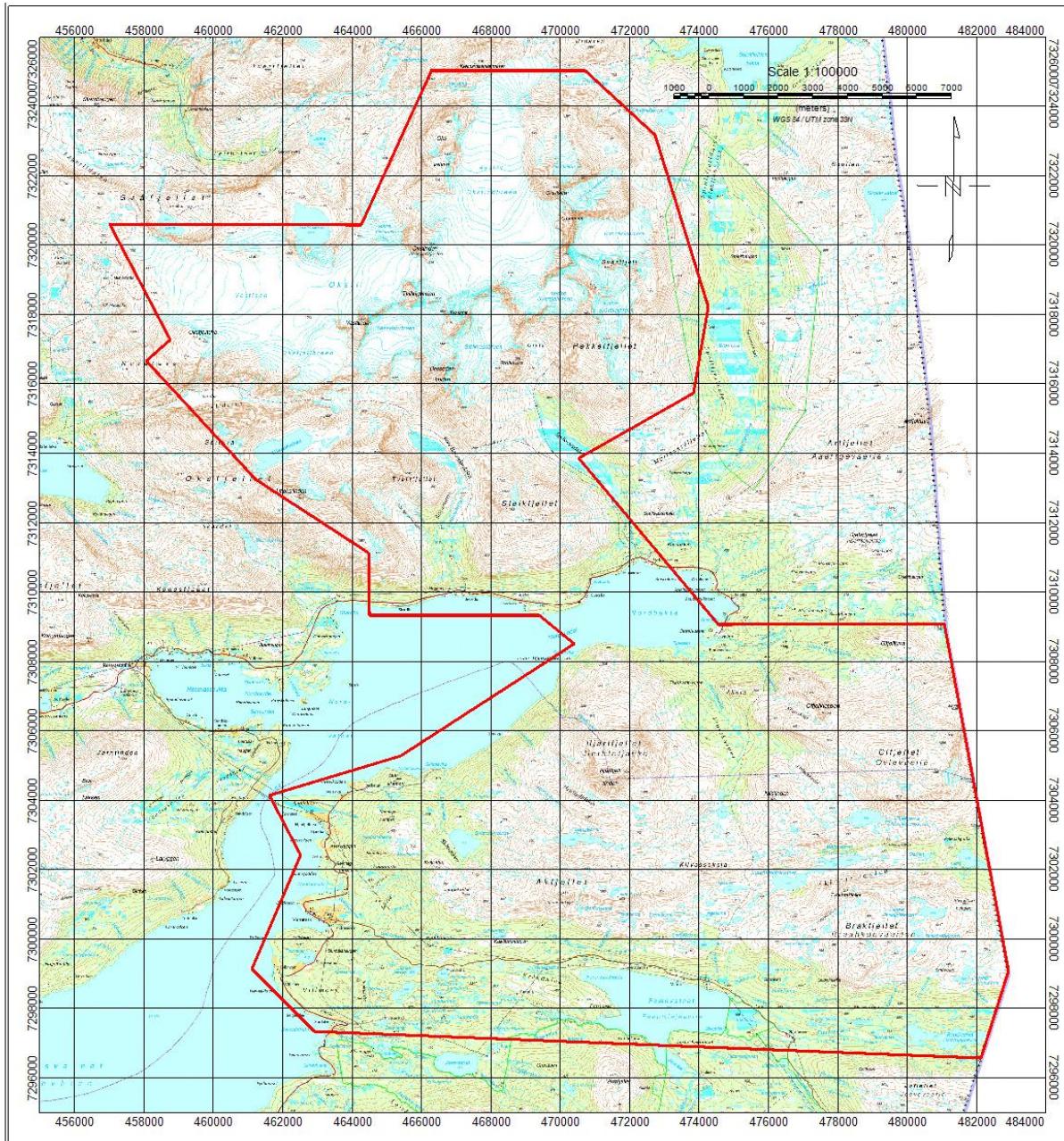


Figure 1: Helicopter survey area in Okstindan and Røsvatnet area.

Table 1. Flight specifications

Survey name	Surveyed lines (km)	Surveyed area (Km ²)	Line direction	Average speed (km/h)
Okstindan 2021	2039	394.5	90	90

The objective of the airborne geophysical survey was to obtain a dense high-resolution magnetic, electromagnetic and radiometric data set over the survey area. These data are required for the enhancement of a general understanding of the regional geology of the area, with adjoining areas covered by recent airborne surveys.

In this regard, the new data can be used to map contacts and structural features within the survey area. It also improves defining the potential of known zones of mineralization, their geological settings, and identifying new areas of interest, as the dataset fills a gap in the high-resolution geophysical surveys of the region.

The survey incorporated the use of a Hummingbird™ 5-frequency electromagnetic (EM) system supplemented by a high-sensitivity cesium magnetometer, gamma-ray spectrometer, and a helicopter mounted radar altimeter. A GPS navigation computer system with flight path indicators ensured accurate positioning of the geophysical data with respect to the World Geodetic System 1984 geodetic datum (WGS-84).

1. SURVEY SPECIFICATIONS

1.1 Airborne survey parameters

NGU used a modified Hummingbird™ EM and magnetic helicopter survey system designed to obtain low level, slow speed, detailed airborne EM and magnetic data (Geotech 1997). The system was supplemented by a Radiation Solutions RSX-5, 1024 channel gamma-ray spectrometer installed between the landing skids under the cabin of the helicopter which is used to map ground concentrations of U, Th and K, and radiation Total Counts.

The airborne survey data were acquired from August 13th to August 27th, 2021 in the Okstindan and Røsvatnet area. A Eurocopter AS350-B3 (LN-OSD) from helicopter company Pegasus Helicopter AS was used during the survey. The survey lines were spaced 200 meters apart, with lines oriented at 90° in UTM zone 33. The magnetic and electromagnetic sensors are housed in a single 7 m long bird, towed 30 meters below the helicopter, and flown at an average elevation of about 65 m above the topographic surface.

Rugged terrain and abrupt changes in topography affected the aircraft pilot's ability to 'drape' the terrain, meaning the average instrumental height was sometimes higher than the standard survey instrumental height, which is defined as 30 meters plus a height of obstacles (trees, power lines etc.) for EM and magnetic sensors.

The ground speed of the helicopter varied from 23 – 128 km/h depending on topography, wind direction and its magnitude. On average the ground speed during measurements is calculated to 90 km/h. Magnetic data were recorded at 0.2 second intervals resulting in approximately 5 meters average point spacing.

EM data were recorded at 0.1 second intervals resulting in data with a sample increment of 2.5 meters along the ground in average. Spectrometry data were recorded every 1 second giving a point spacing of approximately 25 meters. The above parameters allow sufficient detail recognition in the data to detect subtle anomalies that may represent mineralization and/or rocks of different lithological and petrophysical composition.

A base magnetometer to monitor diurnal variations in the magnetic field was located at the base at Vesterli Skistadion, west of Korgen in Hemnes, in August. The GEM GSM-19 base station magnetometer data were recorded once every 3 seconds. The CPU clock of the base magnetometer and the helicopter magnetometer were both synchronized to UTC (Universal Time Coordinates) through the built-in GPS receiver to allow correction of diurnals.

Navigation system uses GPS/GLONASS satellite tracking systems to provide real-time WGS-84 coordinate locations for every second. The accuracy achieved with no differential corrections is reported to be ± 5 meters in the horizontal directions. The GPS receiver antenna was mounted internally inside the canopy of the helicopter.

For quality control, the electromagnetic, magnetic, and radiometric, altitude and navigation data were monitored on four separate windows in the operator's display during flight while they were recorded in three data ASCII streams to the PC hard disk drive. Spectrometry data were also recorded to an internal hard drive of the spectrometer. The data files were transferred to the field workstation via USB flash drive. The raw data files were backed up onto USB flash drive in the field.

1.2 Airborne survey instrumentation

Instrument specifications are given in Table 2. Frequencies and coil configuration for the Hummingbird EM system is given in Table 3.

Table 2. Instrument Specifications

Instrument	Producer/Model	Accuracy / Sensitivity	Sampling frequency / interval
Magnetometer	Scintrex Cs-2	<2.5nT throughout range / 0.0006nT $\sqrt{\text{Hz rms}}$	5 Hz
Base magnetometer	GEM GSM-19	0.1 nT	3 s
Electromagnetic	Geotech Hummingbird	1 – 2 ppm	10 Hz
Gamma spectrometer	Radiation Solutions RSX-5	1024 channels, 16 liters down, 4 liters up	1 Hz
Radar altimeter	Bendix/King KRA 10A	± 5 ft 40 – 100 feet $\pm 5\%$ 100 – 500 feet $\pm 7\%$ 500 – 2500 feet	1 Hz
Pressure/temperature	Honeywell PPT	$\pm 0.03\%$ FS	1 Hz
Navigation	Topcon GPS-receiver	± 5 meters	1 Hz
Acquisition system	NGU custom software		



Figure 2: Pegasus helicopter with pilot and NGU Hummingbird EM system

Table 3. Hummingbird EM system, frequency, and coil configurations

Coils	Frequency	Orientation	Separation
A	7700 Hz	Coaxial	6.30 m
B	6600 Hz	Coplanar	6.30 m
C	980 Hz	Coaxial	6.025 m
D	880 Hz	Coplanar	6.025 m
E	34133 Hz	Coplanar	4.90 m

1.3 Airborne Survey Logistics Summary

A summary of the survey specifications is shown in Table 4.

Table 4. Survey specifications summary

Parameter	Specifications
Traverse (survey) line spacing	200 meters
Traverse line direction	E-W (90°)
Nominal aircraft ground speed	23 - 128 km/h
Average aircraft ground speed	90 km/h
Average sensor terrain clearance Mag	65 meters
Average sensor terrain clearance Rad	93 meters
Sampling rates:	
Magnetometer	0.2 seconds
Electromagnetic system	0.1 seconds
Spectrometer, GPS, altimeter	1.0 second
Base Magnetometer	3.0 seconds

2. DATA PROCESSING AND PRESENTATION

The survey planning was done by Frode Ofstad. The field data quality control was performed by Marie-Andrée Dumais. Data acquisition was done by Frida Mathayo Mrope and Tom Kristiansen. The magnetic, spectrometry were processed by Alexandros Stampolidis and resistivity data were processed by Tom Kristiansen at NGU. The ASCII data files were loaded into three separate Oasis Montaj databases. All three datasets were processed consequently according to processing flow charts shown in Appendix A1, A2 and A3.

2.1 Total Field Magnetic Data

At the first stage the raw magnetic data were visually inspected, and spikes were removed manually. Non-linear filter was also applied to airborne raw data to eliminate short-period spikes. Typically, several corrections must be applied to magnetic data before gridding - heading correction, lag correction and diurnal correction.

Diurnal corrections

The temporal fluctuations in the magnetic field of the earth affect the total magnetic field readings recorded during the airborne survey. This is commonly referred to as the magnetic diurnal variation. These fluctuations can be effectively removed from the airborne magnetic dataset by using a stationary reference magnetometer that records the magnetic field of the earth simultaneously with the airborne sensor at given short time interval.

Diurnal variation data were inspected for spikes, and spikes were removed manually if necessary. Magnetic diurnals that were recorded on the base station magnetometer were within the standard NGU specifications during the entire survey (Rønning 2013).

Diurnal variations were measured with GEM GSM-19 magnetometer. The base station computer clock was continuously synchronized with GPS clock. The recorded data are merged with the airborne data and the diurnal correction is applied according to equation (1).

$$\mathbf{B}_{Tc} = \mathbf{B}_T + (\bar{\mathbf{B}}_B - \mathbf{B}_B), \quad (1)$$

Where:

\mathbf{B}_{Tc} = Corrected airborne total field readings

\mathbf{B}_T = Airborne total field readings

$\bar{\mathbf{B}}_B$ = Average datum base level

\mathbf{B}_B = Base station readings

The average datum base level ($\bar{\mathbf{B}}_B$) was set to 52868 nT for the survey flown from the base at Vesterli Skistation in Korgen, Hemnes.

Corrections for lag and heading

Neither a lag nor a cloverleaf test was performed before the survey. According to previous reports the lag between logged magnetic data and the corresponding navigational data was 1-2 fids. These values were observed to have a negligible effect on the processed results. A heading error for a towed system is usually either very small or non-existent. No lag and heading corrections were applied.

Magnetic data processing, gridding, and presentation

The total field magnetic anomaly data (\mathbf{B}_{TA}) were calculated from the diurnal corrected data (\mathbf{B}_{Tc}) after subtracting the IGRF for the surveyed area calculated for the data period (eq.2)

$$\mathbf{B}_{TA} = \mathbf{B}_{Tc} - IGRF \quad (2)$$

IGRF 2020 model was employed in these calculations, to ensure that the Okstindan data set would match the previously processed and earlier published surrounding data sets.

The total field anomaly data were split into lines and then were gridded using a minimum curvature method with a grid cell size of 50 meters. This cell size is exactly one quarter of the 200 meters average line spacing. To remove small line-to-line levelling errors that were detected on the gridded magnetic anomaly data, micro-levelling technique was applied on the flight line based magnetic database. Finally, the micro-leveled data were gridded using minimum curvature method with 50 meters grid cell size.

The processing steps of magnetic data presented so far, were performed on point basis. The following steps are performed on grid basis.

The horizontal and vertical gradient along with the tilt derivative of the total magnetic anomaly were calculated from the stitched micro-leveled total magnetic anomaly grid. The magnitude of the horizontal gradient was calculated according to equation (3)

$$HG = \sqrt{\left(\frac{\partial(B_{TA})}{\partial x}\right)^2 + \left(\frac{\partial(B_{TA})}{\partial y}\right)^2} \quad (3)$$

where \mathbf{B}_{TA} is the micro-leveled total field anomaly field. The vertical gradient (VG) was calculated by applying a vertical derivative convolution filter to the micro-leveled \mathbf{B}_{TA} field. The tilt derivative (TD) was calculated according to the equation (4)

$$TD = \tan^{-1}\left(\frac{VG}{HG}\right) \quad (4)$$

A 3x3 convolution filter was applied to smooth the resulted magnetic grids.

The results are presented in a series of colored shaded relief maps (1:100000). The maps are:

- Total field magnetic anomaly
- Horizontal gradient of total magnetic anomaly
- Vertical gradient of total magnetic anomaly
- Tilt derivative (or Tilt angle) of the total magnetic anomaly

These maps are representative of the distribution of magnetization over the surveyed areas. The list of the produced maps is shown in Table 6.

2.2 Electromagnetic data

The EM system transmits five fixed frequencies and records an in-phase (IP) and a quadrature (Q) response for each of the five coil sets of the electromagnetic system. The received signals are processed and used for calculation of apparent resistivity.

IP and Q data were filtered with 15 fiducial non-linear filter to eliminate spherical spikes, which were represented as irregular noise of large amplitude in records and high frequency noise of bird electronics. Then, a 20-fiducial low-pass filter was applied to suppress instrumental and cultural noise. These filters were not able to suppress all the noise. Also, shifts in IP and Q data, with amplitude of 5-10 ppm, was observed in some flights. These shifts were edited manually where possible.

To remove the effects of instrument drift caused by gradual temperature variations in the transmitting and receiving circuits, background responses are recorded during each flight. To obtain a background level, the bird is raised to an altitude of at least 1000 ft above the topographic surface so that no electromagnetic responses from the ground are present in the recorded traces.

The EM traces observed at this altitude correspond to a background (zero) level of the system. If these background levels are recorded at 20-30 minutes interval, then the linear drift of the system can be removed on a flight-by-flight basis, before any further processing is carried out. Geosoft HEM module was used for applying drift correction. Residual instrumental drift, usually small, but non-linear, was manually removed manually on a line-to-line basis.

A 5 ppm amplitude-limited, 120 seconds low-pass (LP) filter was applied to the in-phase and quadrature signals. The LP filter result was then subtracted from each component before the inversion calculations. The LP filtering of the EM data produces more uniform data and clearer anomalies with fewer linear artifacts in the grids.

When final levelling of the EM data was complete, apparent resistivity was calculated from in-phase and quadrature EM components using a homogeneous half space model of the earth (Geosoft HEM module) for 6600, 7000, 980, 880 and 34133 Hz. using a threshold value of 3 ppm, starting value of 500 ohm-m, and fractional error 1%.

Electromagnetic field decays rapidly with the distance (height of the sensors) – as z^{-2} – z^{-5} depending on the shape of the conductors and, at certain height, signals from the ground sources become comparable with instrumental noise. Levelling errors or precision of levelling can sometimes lead to appearance of artificial resistivity anomalies when data were collected at high instrumental altitude.

Application of threshold allows excluding such data from an apparent resistivity calculation, though not completely. It is particularly noticeable in low frequencies datasets. Resistivity data were visually inspected; artificial anomalies associated with high altitude measurements were manually removed.

Data recorded at a height larger than 150 meters were considered as non-reliable and removed from presentation. The apparent resistivity was gridded with a cell size of 50 meters. Power lines strongly affected low frequency data – 880 and 980 Hz

frequencies, and the most prominent noise from large power lines were masked in the final grids.

2.3 Radiometric data

Airborne gamma-ray spectrometry measures the abundance of Potassium (K), Thorium (eTh), and Uranium (eU) in rocks and weathered materials by detecting gamma-rays emitted due to the natural radioelement decay of these elements. The data analysis method is based on the IAEA recommended method for U, Th and K (International Atomic Energy Agency, 1991; 2003). A short description of the individual processing steps of that methodology as adopted by NGU is given below.

Energy windows

The Gamma-ray spectra were initially reduced into standard energy windows corresponding to the individual radio-nuclides K, U and Th. Figure 3 shows an example of a Gamma-ray spectrum and the corresponding energy windows and radioisotopes (with peak energy in MeV) responsible for the radiation.

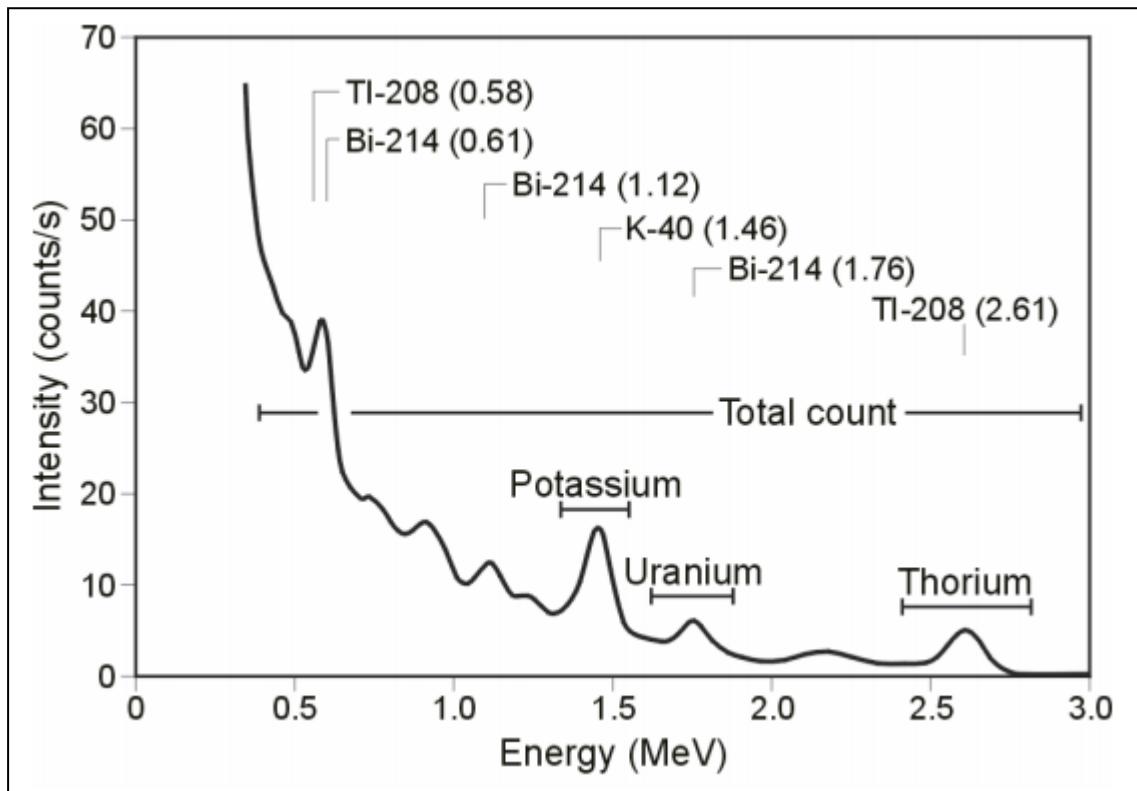


Figure 3: Gamma-ray spectrum with K, Th, U and Total Count windows.

Table 5. Specified channel windows for the 1024 RSX-5 system.

Gamma-ray spectrum	Cosmic	Total count	K	U	Th
Down	1022	135-935	455-522	552-618	802-935
Up	1022			552-618	
Energy windows (MeV)	>3.07	0.4-2.8	1.36-1.56	1.65-1.85	2.4-2.8

The RSX-5 is a 1024 channel system with four downward and one upward looking detector, which means that the actual Gamma-ray spectrum is divided into 1024 channels. The first channel is reserved for the “Live Time” and the last for the Cosmic rays. Table 5 shows the windows that were used for the reduction of the spectrum.

Live Time correction

The data were corrected for live time. “Live time” is an expression of the relative length of time the instrument was able to register new pulses per sample interval. On the other hand, “dead time” is an expression of the relative length of time the system was unable to register new pulses per sample interval. The relation between “dead time” and “live time” is given by the equation (5)

$$\text{“Live time”} = \text{“Real time”} - \text{“Dead time”} \quad (5)$$

where the “real time” or “acquisition time” is the elapsed time over which the spectrum is accumulated (about 1 second).

The live time correction is applied to the total count, Potassium, Uranium, Thorium, upward Uranium, and cosmic channels. The formula used to apply the correction is as follows:

$$C_{LT} = C_{RAW} \cdot \frac{\text{Acquisition Time}}{\text{Live Time}} \quad (6)$$

where C_{LT} is the live time corrected window in counts per second, C_{RAW} is the raw window data in counts per second, while Acquisition Time and Live Time are in microseconds.

Cosmic and aircraft correction

Background radiation resulting from cosmic rays and aircraft contamination was removed from the total count, Potassium, Uranium, Thorium, upward Uranium window using the following formula:

$$C_{CA} = C_{LT} - (a_c + b_c \cdot C_{Cos}) \quad (7)$$

where C_{CA} is the cosmic and aircraft corrected window, C_{LT} is the live time corrected window, a_c is the aircraft background for this window, b_c is the cosmic stripping coefficient for this window and C_{Cos} is the low pass filtered cosmic window. Cosmic and aircraft background coefficients are determined by high altitude calibration flights.

Radon correction

The upward detector method, as discussed in IAEA (1991), was applied to remove the effects of the atmospheric radon in the air below and around the helicopter. Using spectrometry data over-water, where there is no contribution from the ground sources, enables the calculation of the coefficients (a_c and b_c) for the linear equations that relate the cosmic corrected counts per second of Uranium window with that of total count, Potassium, Thorium and Uranium upward window over water. Data over-land were used in conjunction with data over-water to calculate the a_1 and a_2 coefficients used in equation (8) for the determination of the Radon component in the downward Uranium window:

$$Radon_U = \frac{U_{upCA} - a_1 \cdot U_{CA} - a_2 \cdot Th_{CA} + a_2 \cdot b_{Th} - b_U}{a_U - a_1 - a_2 \cdot a_{Th}} \quad (8)$$

where $Radon_U$ is the Radon component in the downward Uranium window, U_{upCA} is the filtered upward uranium, U_{CA} is the filtered Uranium, Th_{CA} is the filtered Thorium, a_1 , a_2 , a_U and a_{Th} are proportional factors and b_U and b_{Th} are constants determined experimentally.

The effects of Radon in the downward Uranium are removed by simply subtracting $Radon_U$ from U_{CA} . The effects of radon in the other windows are removed using the following formula:

$$C_{RC} = C_{CA} - (a_C \cdot Radon_U + b_C) \quad (9)$$

where C_{RC} is the Radon corrected window, C_{CA} is the cosmic and aircraft corrected window, $Radon_U$ is the Radon component in the downward uranium window, a_C is the proportionality factor and b_C is the constant determined experimentally for this window from over-water data.

Compton stripping

Radon corrected Potassium, Uranium and Thorium windows are subjected to spectral overlap correction. Compton scattered gamma rays in the radio-nuclides energy windows were corrected by window stripping using Compton stripping coefficients determined from measurements on calibrations pads (Grasty et al, 1991) at the Geological Survey of Norway in Trondheim (see values in Appendix A2).

The stripping corrections are given by the following formulas:

$$A_l = 1 - (g \cdot \gamma) - (a \cdot \alpha) + (a \cdot g \cdot \beta) - (b \cdot \beta) + (b \cdot \alpha \cdot \gamma) \quad (10)$$

$$U_{ST} = \frac{Th_{RC} \cdot ((g \cdot \beta) - \alpha) + U_{RC} \cdot (1 - b \cdot \beta) + K_{RC} \cdot ((b \cdot \alpha) - g)}{A_l} \quad (11)$$

$$Th_{ST} = \frac{Th_{RC} \cdot (1 - (g \cdot \gamma)) + U_{RC} \cdot (b \cdot \gamma - a) + K_{RC} \cdot ((a \cdot g) - b)}{A_l} \quad (12)$$

$$K_{ST} = \frac{Th_{RC} \cdot ((\alpha \cdot \gamma) - \beta) + U_{RC} \cdot ((a \cdot \beta) - \gamma) + K_{RC} \cdot (1 - (a \cdot \alpha))}{A_l} \quad (13)$$

where U_{RC} , Th_{RC} , K_{RC} are the Radon corrected Uranium, Thorium and Potassium and a , b , g , α , β , γ are Compton stripping coefficients, determined from measurements on the calibration pads at NGU. U_{ST} , Th_{ST} and K_{ST} are stripped values of Uranium, Thorium and Potassium.

Reduction to Standard Temperature and Pressure

The radar altimeter data were converted to effective height (H_{STP}) using the acquired temperature and pressure data, according to the expression:

$$H_{STP} = H \cdot \frac{273.15}{T + 273.15} \cdot \frac{P}{1013.25} \quad (14)$$

where H is the smoothed observed radar altitude in meters, T is the measured air temperature in degrees Celsius and P is the measured barometric pressure in millibars.

Height correction

Variations caused by changes in the aircraft altitude relative to the ground were corrected to a nominal height of 60 meters. Data recorded at a height larger than 150 meters were considered as non-reliable and removed from processing. Total count, Uranium, Thorium and Potassium stripped windows were subjected to height correction according to the equation:

$$C_{60m} = C_{ST} \cdot e^{C_{ht} \cdot (60 - H_{STP})} \quad (15)$$

where C_{ST} is the stripped corrected window, C_{ht} is the height attenuation factor for that window, determined from a calibration flight, and H_{STP} is the effective height.

Conversion to ground concentrations

Finally, corrected count rates were converted to effective ground element concentrations using calibration values derived from calibration pads (Grasty et al, 1991) at the Geological Survey of Norway in Trondheim (see values in Appendix A2). The corrected data provide an estimate of the apparent surface concentrations of Potassium, Uranium and Thorium (K, eU and eTh). Potassium concentration is expressed as a percentage, equivalent Uranium and Thorium as parts per million (ppm). Uranium and Thorium are described as “equivalent” since their presence is inferred from gamma-ray radiation from daughter elements (^{214}Bi for Uranium, ^{208}Tl for Thorium). The concentration of the elements is calculated according to the following expressions:

$$C_{CONC} = C_{60m} / C_{SENS_60m} \quad (16)$$

where C_{60m} is the height corrected window, C_{SENS_60m} is experimentally determined sensitivity reduced to the nominal height (60m).

Spectrometry data gridding and presentation

Gamma-rays from Potassium, Thorium and Uranium emanate from the uppermost 30 to 40 centimeters of soil and rock in the crust (Minty, 1997). Variations in the concentrations of these radioactive elements are largely related to changes in the mineralogy and geochemistry of the Earth’s surface.

The spectrometry data were stored in a database and the ground concentrations were calculated following the processing steps. A list of the parameters used in these steps is given in Appendix A3.

Then the data were split in lines and ground concentrations of the three main natural radio-elements Potassium, Thorium and Uranium and total gamma-ray flux (total count) were gridded using a minimum curvature method with a grid cell size of 50 meters. To remove small line-to-line levelling errors appeared on those grids, the data were micro-leveled as in the case of the magnetic data, and re-gridded with the same grid cell size.

Quality of the radiometric data was within standard NGU specifications (Rønning 2013). For further reading regarding standard processing of airborne radiometric data, we recommend the publications from Minty et al. (1997).

A 3x3 convolution filter was applied to smooth the concentration grids. A list of the produced maps is shown in Table 6.

3. PRODUCTS

Processed digital data from the survey are presented as:

1. Geosoft XYZ files: Okstindan_Mag.xyz, Okstindan_EM.xyz, Okstindan_Rad.xyz,
Coloured maps at the scale 1:100000 available from NGU on request.
2. Grid-files in Geo-TIFF format

Table 6. Maps in scale 1:100000, available from NGU on request.

Map #	Name
2022.012-00	Survey Flight Path
2022.012-01	Total magnetic field
2022.012-02	Magnetic Horizontal Gradient
2022.012-03	Magnetic Vertical Gradient
2022.012-04	Magnetic Tilt Derivative
2022.012-05	Apparent resistivity, Frequency 7000 Hz, coaxial coils
2022.012-06	Apparent resistivity, Frequency 6600 Hz, coplanar coils
2022.012-07	Apparent resistivity, Frequency 980 Hz, coaxial coils
2022.012-08	Apparent resistivity, Frequency 880 Hz, coplanar coils
2022.012-09	Apparent resistivity, Frequency 34133 Hz, coplanar coils
2022.012-10	Radiometric Total counts
2022.012-11	Potassium ground concentration
2022.012-12	Uranium ground concentration
2022.012-13	Thorium ground concentration
2022.012-14	Radiometric Ternary Image

Downscaled images of the maps are shown in figures 4 to 18.

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Appendix A1: Flow chart of magnetic processing

Meaning of parameters is described in the referenced literature.

Processing flow:

- Creation of database and quality control.
- Visual inspection of airborne data and manual spike removal
- Import of diurnal data from base magnetometer database.
- Correction of data for diurnal variation.
- IGRF 2020 removed.
- Splitting flight data by lines
- Gridding
- Microlevelling
- 3x3 convolution filter
- Microlevelling of Magnetic data using Geosoft menu and smoothening by a convolution filtering.

Microlevelling parameters	Value
De-corrugation cutoff wavelength (m)	600
Cell size for gridding (m)	50
Amplitude limit	5
Naudy (1968) Filter length (m)	1000

Appendix A2: Flow chart of EM processing

Meaning of parameters is described in the referenced literature.

Processing flow:

- Automated leveling using Geosoft HEM module
- Filtering of in-phase and quadrature channels with non-linear and low-pass filters
- Quality control and visual inspection of data.
- Manual removal of remaining part of instrumental drift
- Additional levelling using low-pass filter to reduce linear noise
- Calculation of an apparent resistivity using in-phase and quadrature channels.
- Splitting flight data by lines
- Gridding

Appendix A3: Flow chart of radiometry processing

Underlined processing stages are not only applied to the K, U and Th window, but also to the total count. Meaning of parameters is described in the referenced literature.

- Airborne and cosmic correction (IAEA, 2003)
Used parameters: determined by high altitude calibration flights (1500-9000 ft) at Randsfjorden in October 2021.

Window	Background	Cosmic
K	6.5274	0.0537
U	4.3312	0.0373
Th	0	0.0694
Uup	1.1423	0.0108
Total counts	71.552	0.936

- Radon correction using upward detector method (IAEA, 2003)

Used parameters determined from survey data over water and land at Okstindan, August 2021:

Coefficient	Value	Coefficient	Value
a_u	0.34615	b_u	0.38077
a_K	1.16484	b_K	0.70989
a_{Th}	0.23077	b_{Th}	0.45385
a_{TC}	19.75824	b_{TC}	0.0
a_1	0.04133445	a_2	0.05053322

- Stripping corrections (IAEA, 2003)

Used parameters determined from measurements on calibrations pads at NGU, May 2021

Coefficient	Value
a	0.048987
b	0
c	0
α	0.302131
β	0.463789
γ	0.795178

- Height correction to a height of 60 m

Parameters determined by high altitude calibration flights (100 – 700 ft). The average values from tests performed at Randsfjorden, 2021 were used. Attenuation factors in 1/m:

Window	Attenuation factor
K	-0.0103
U	-0.0093
Th	-0.0085
TC	-0.0088

- Converting counts at 60 m heights to element concentration on the ground

Used parameters determined from measurements on calibrations pads at NGU, May 2021

Window	Sensitivity
K (%/count)	0.00731
U (ppm/count)	0.08489
Th (ppm/count)	0.15411

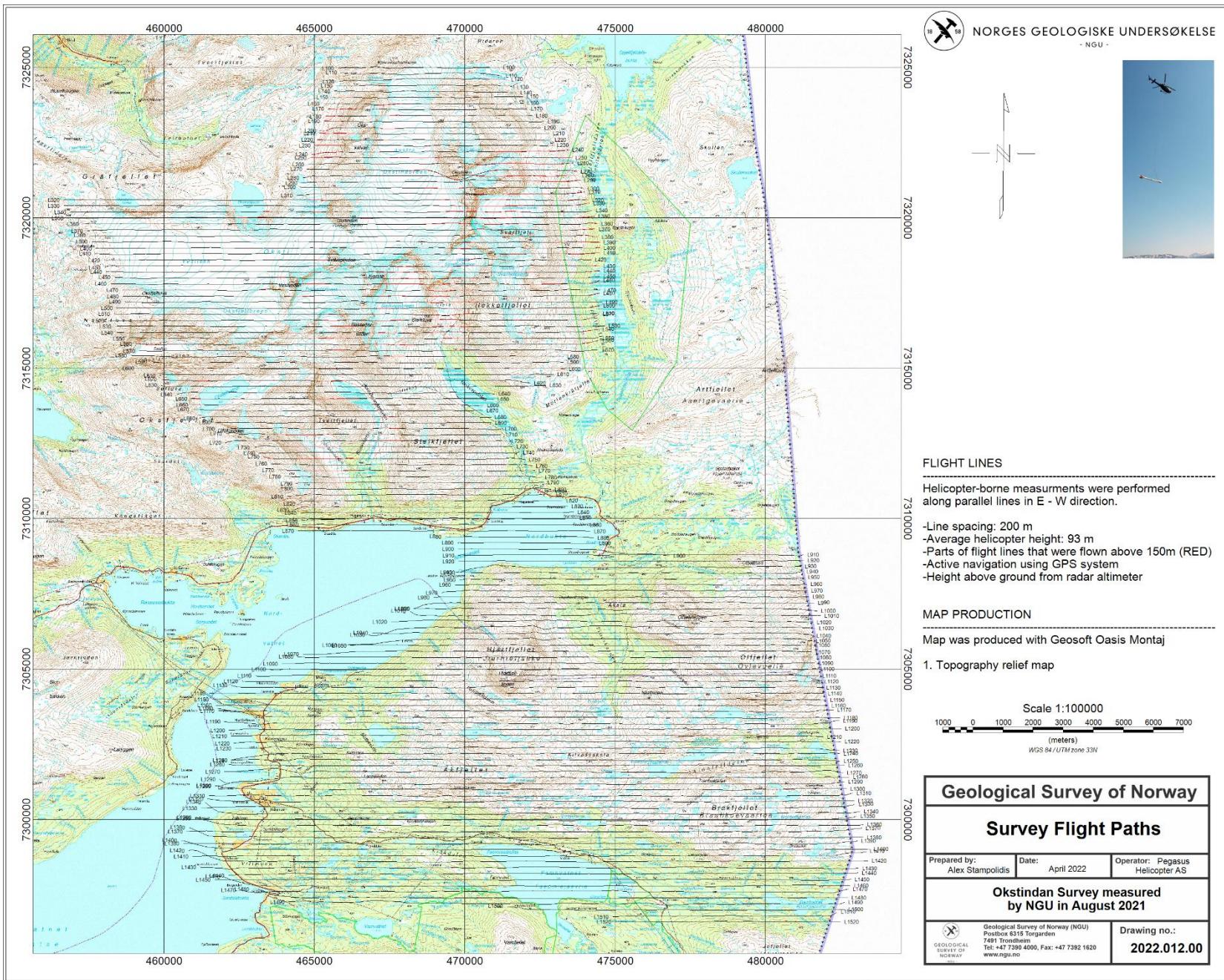


Figure 4: Okstindan and Røsvatnet survey area with flight path. The red segments were flown above 150m.

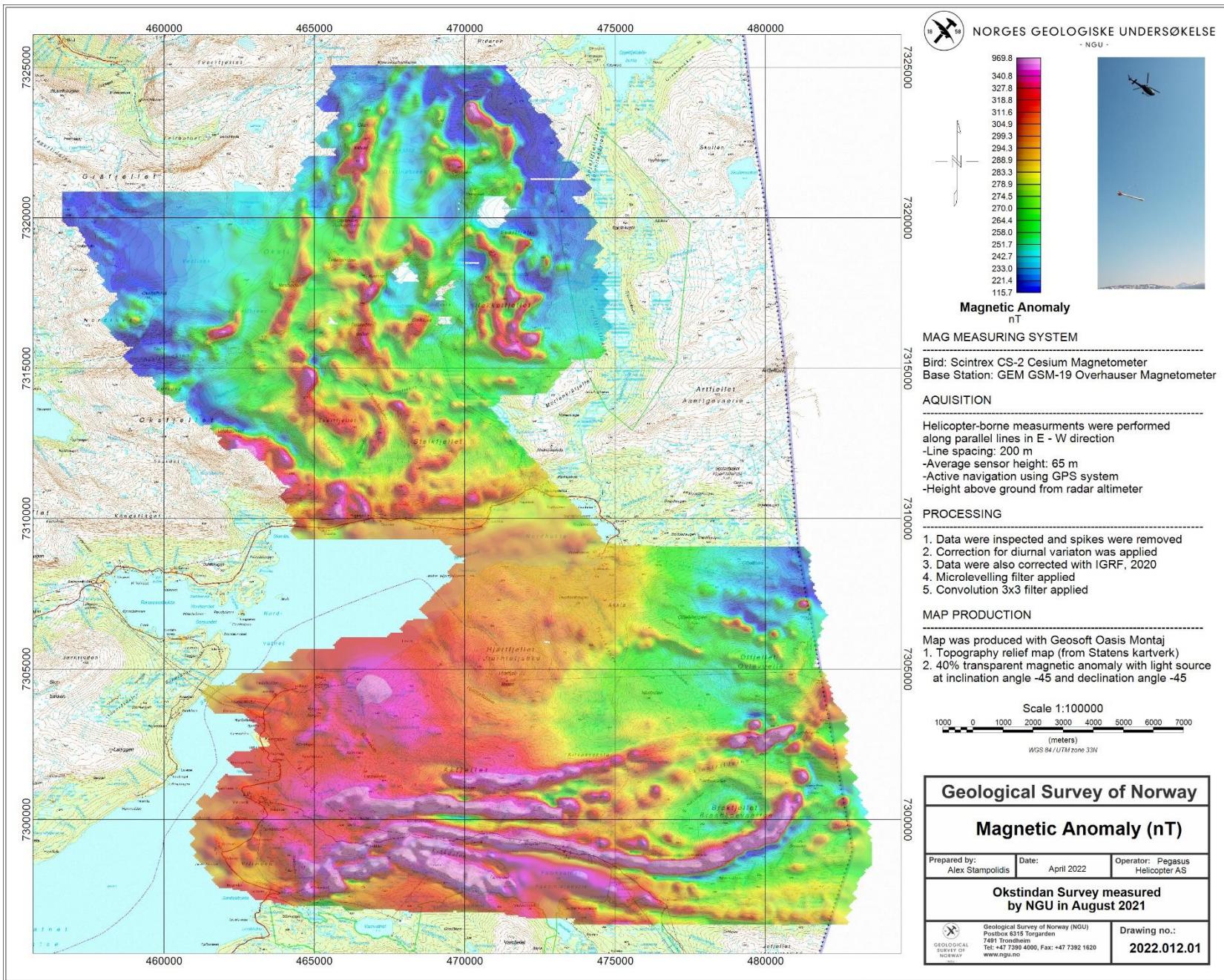


Figure 5: Total Magnetic Field

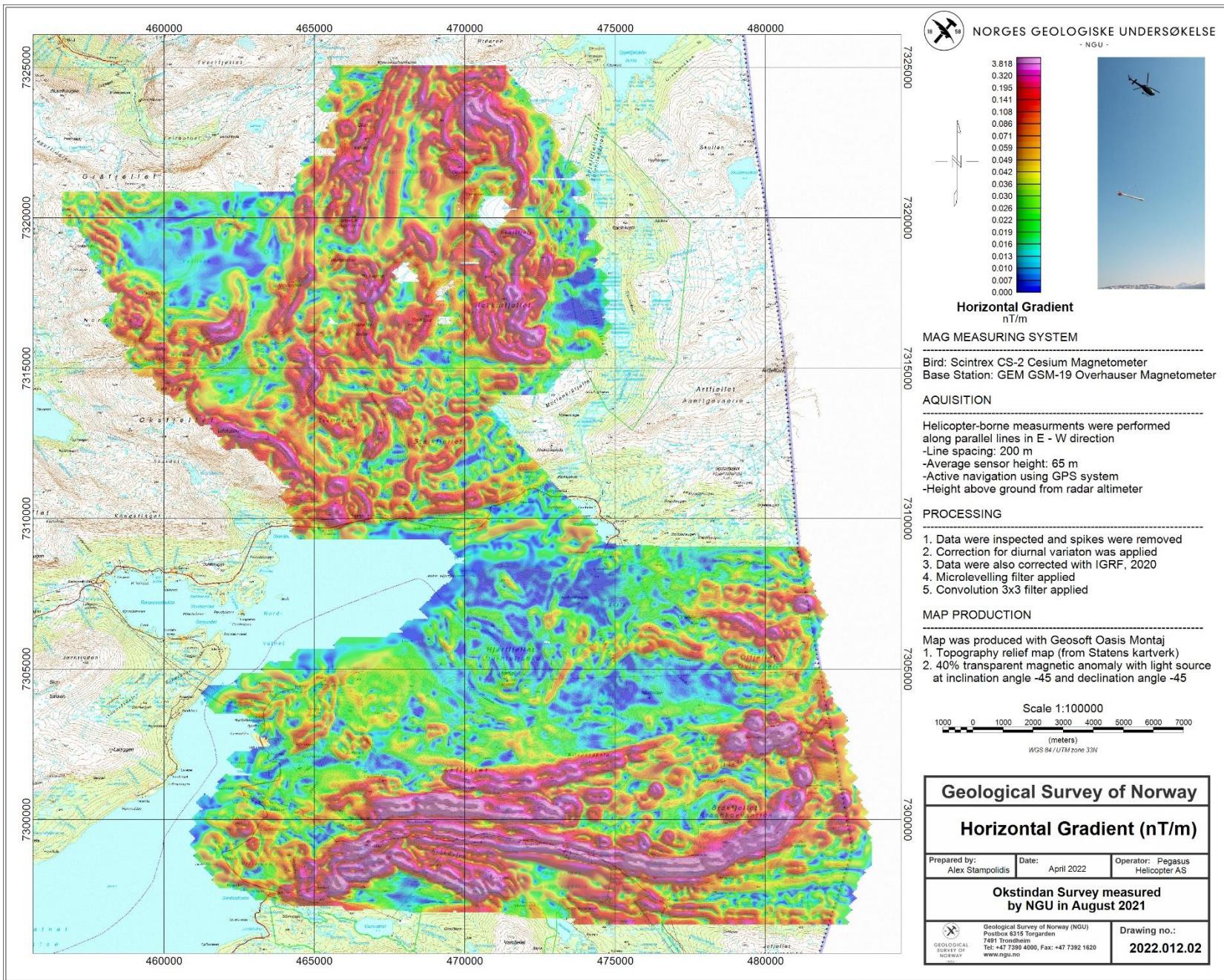


Figure 6: Magnetic Horizontal Gradient

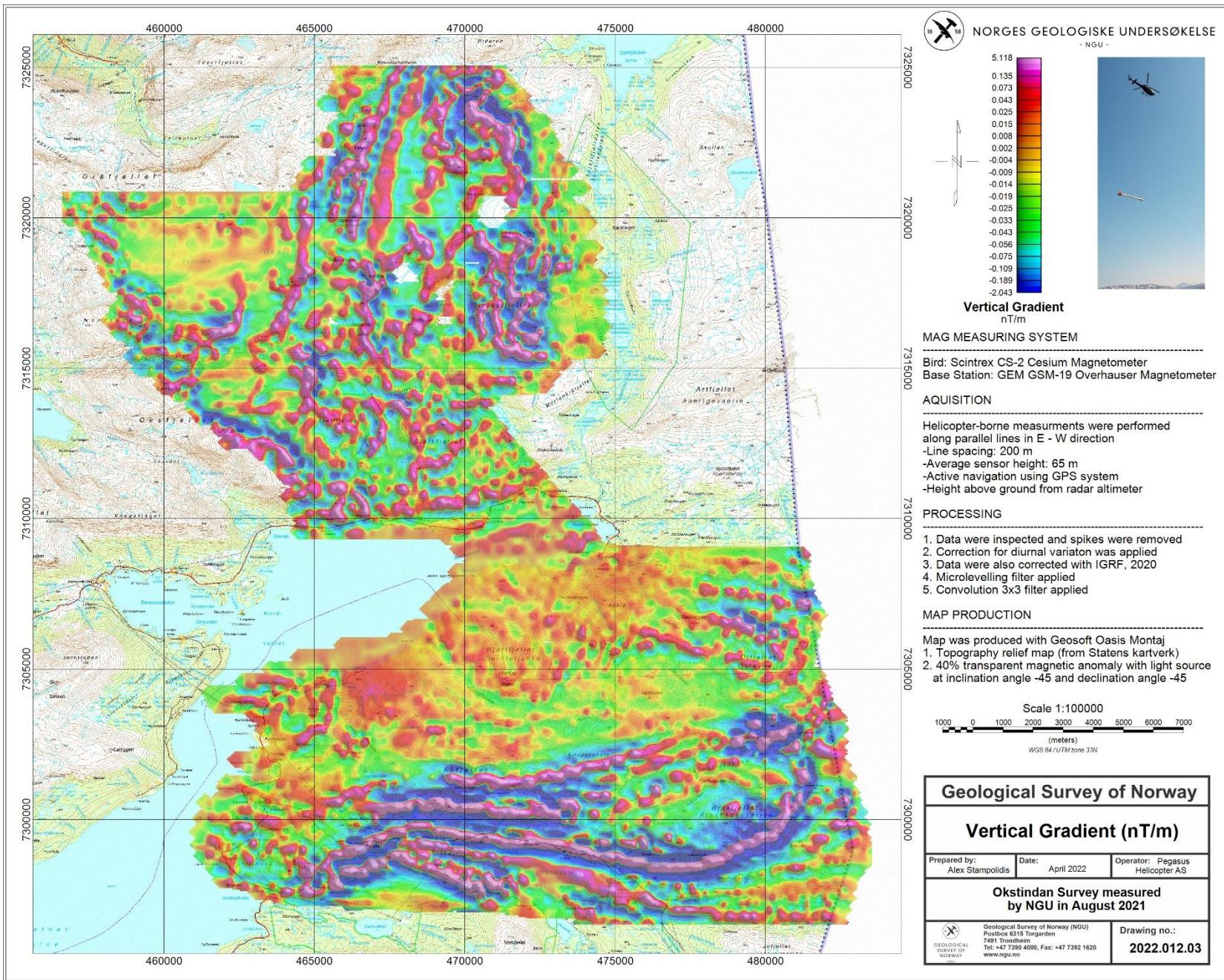


Figure 7: Magnetic Vertical Gradient

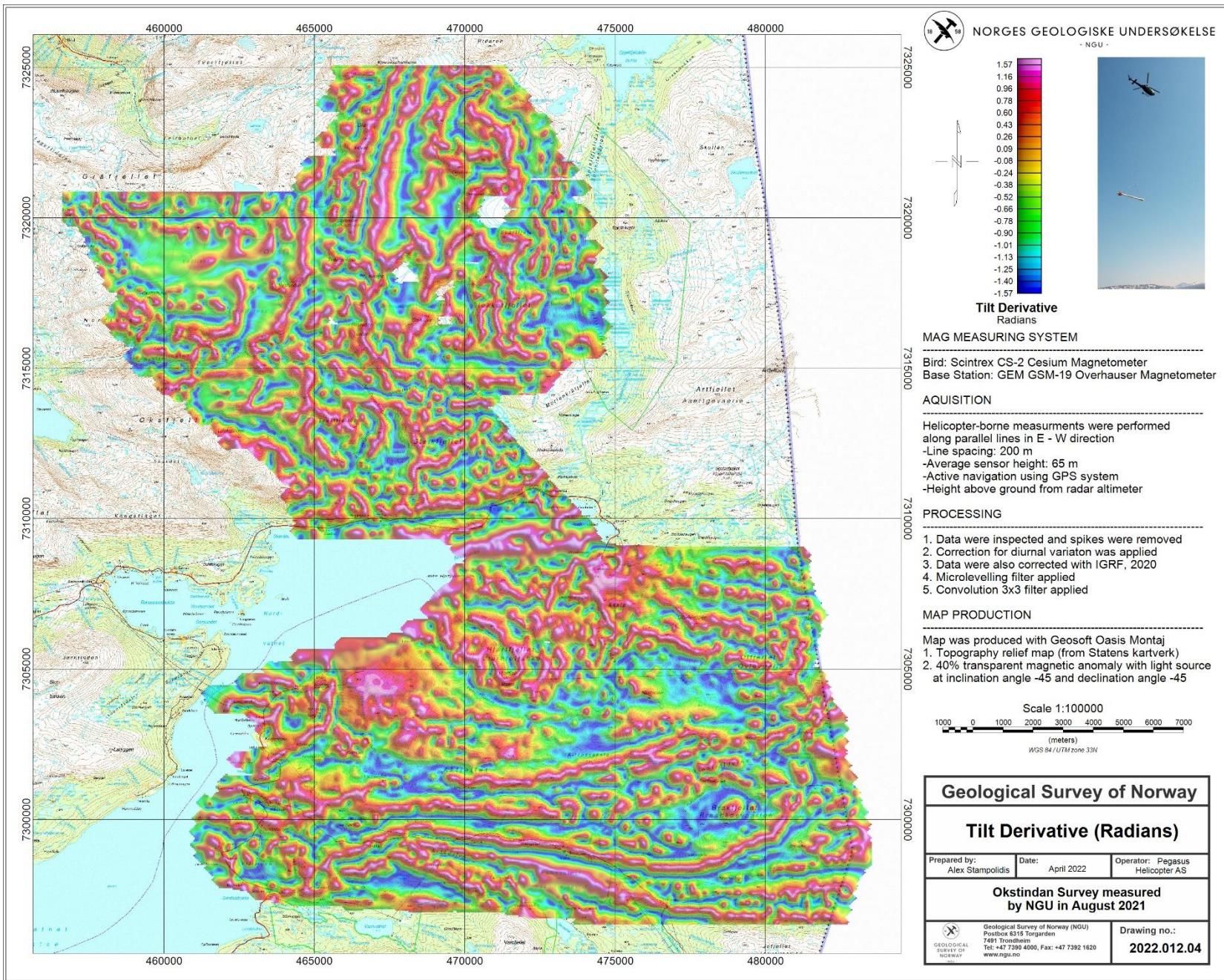


Figure 8: Magnetic Tilt Derivative

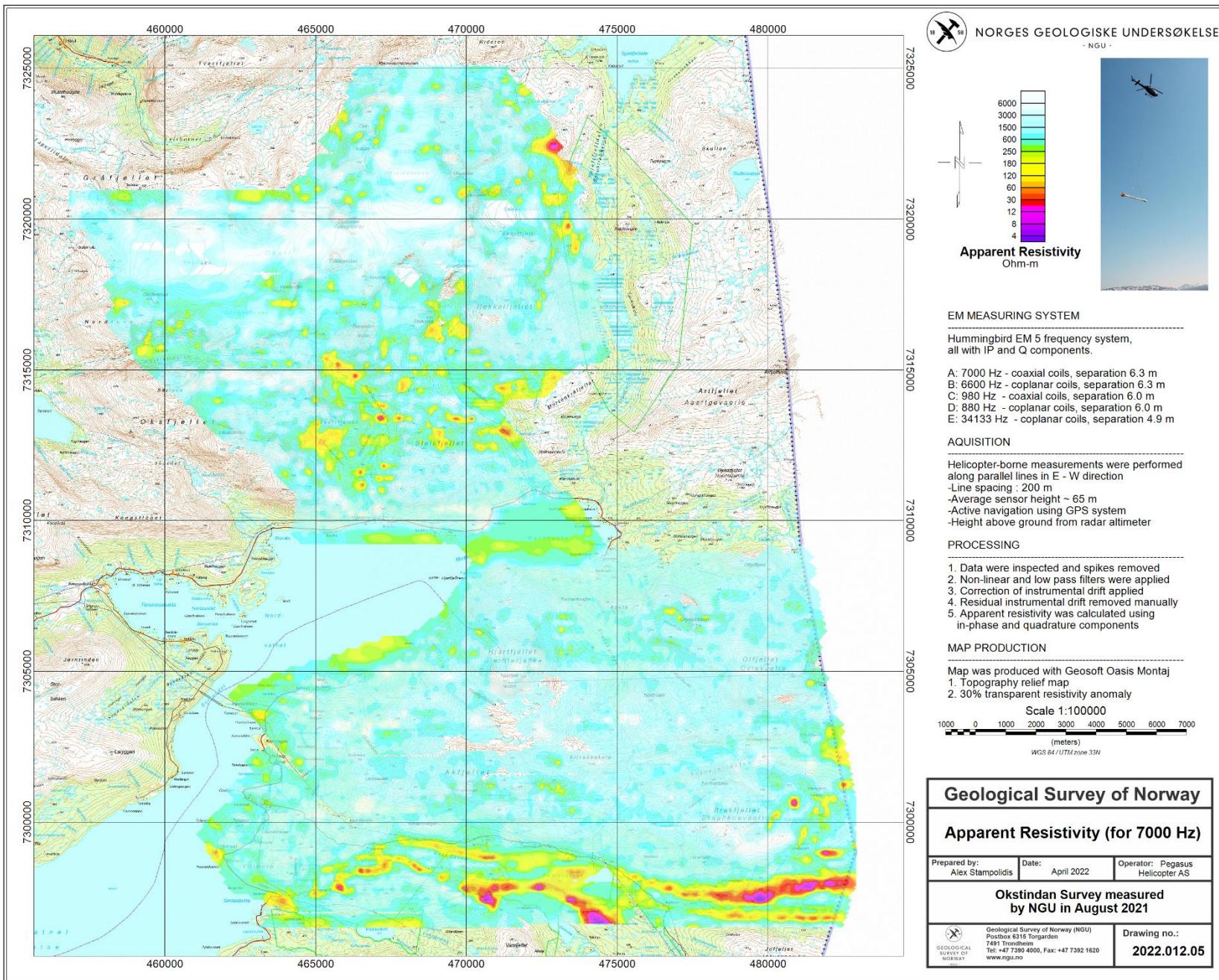


Figure 9: Apparent resistivity. Frequency 7000 Hz, Coaxial coils

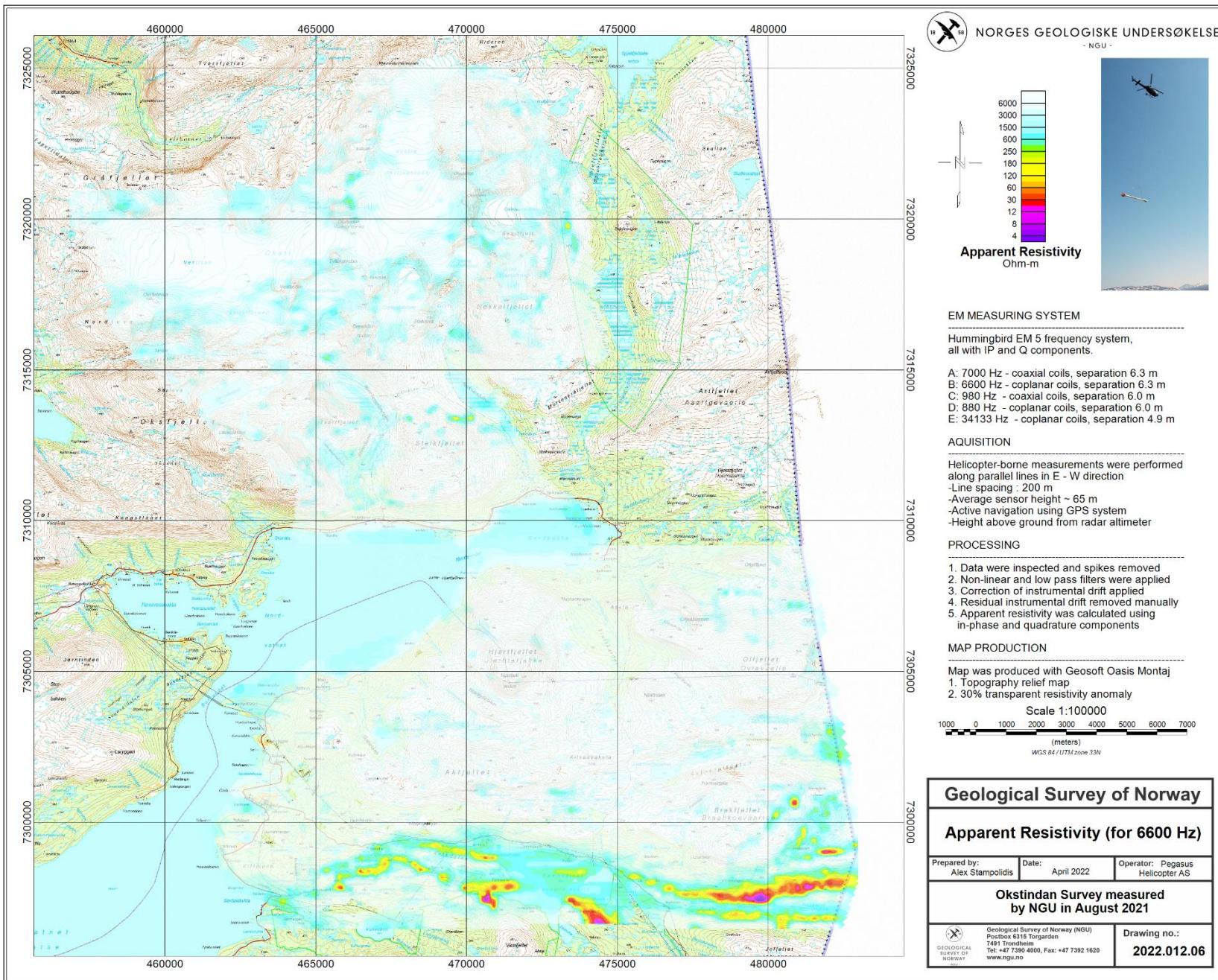


Figure 10: Apparent resistivity. Frequency 6600 Hz, Coplanar coils

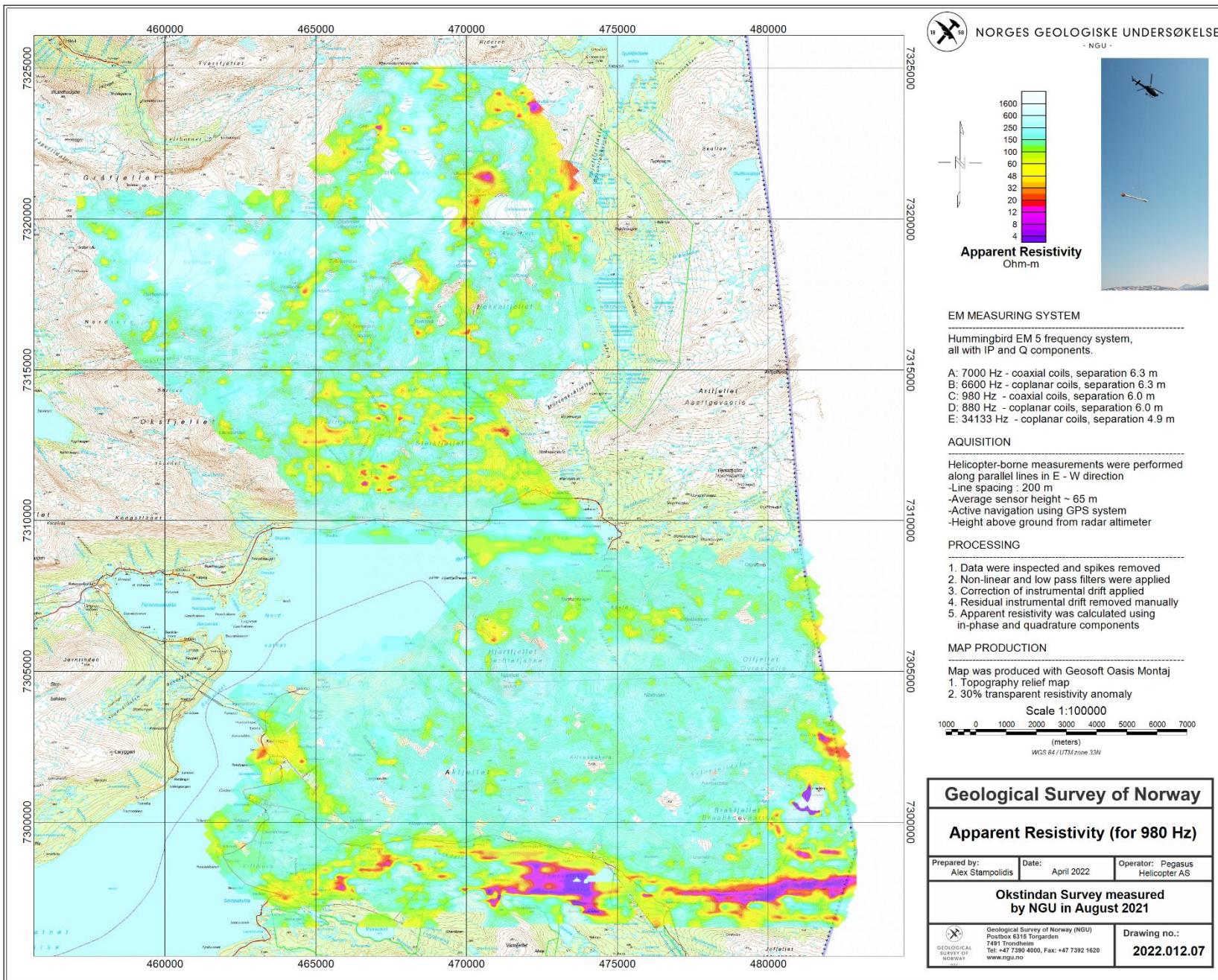


Figure 11: Apparent resistivity. Frequency 980 Hz, Coaxial coils

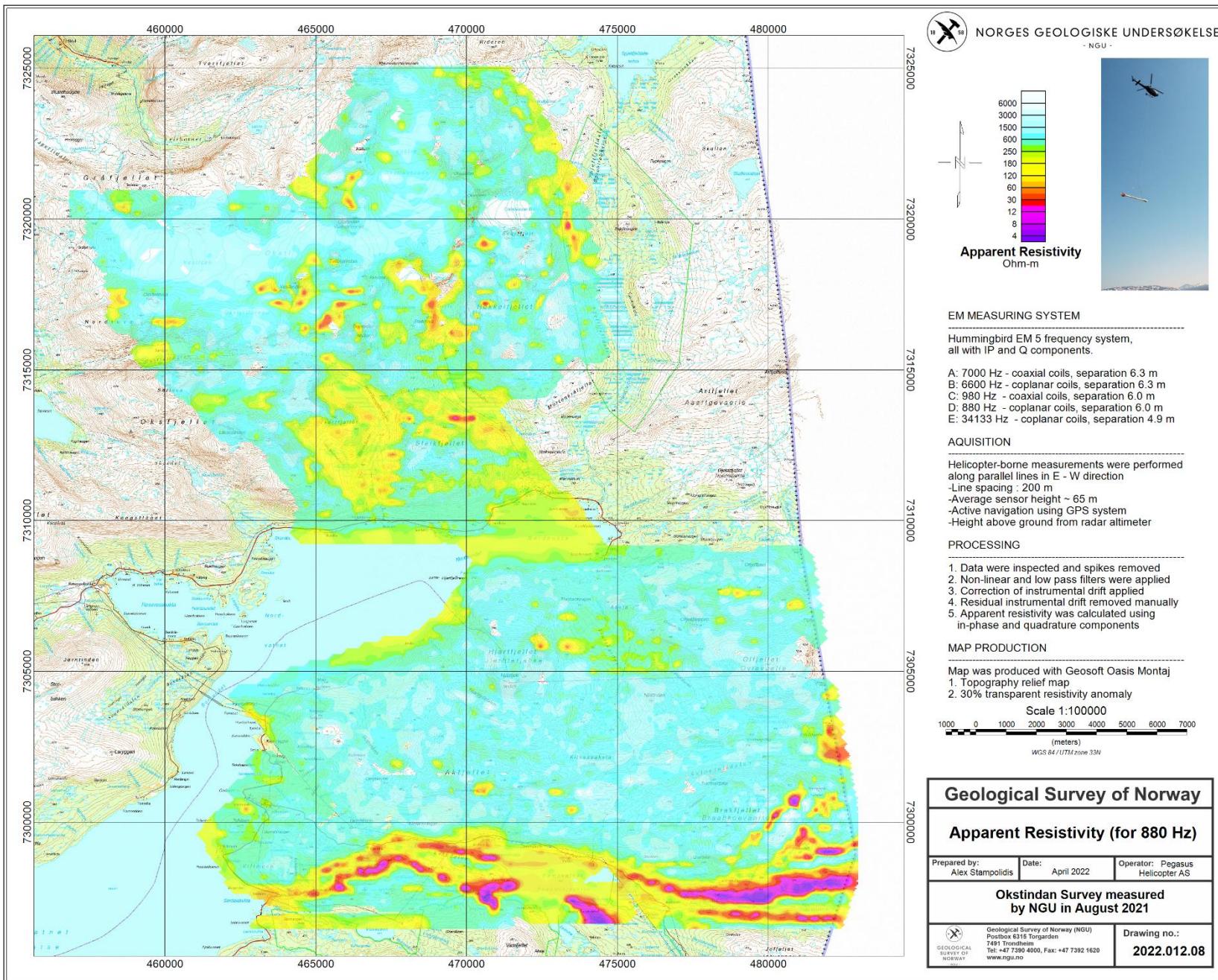


Figure 12: Apparent resistivity. Frequency 880 Hz, Coplanar coils

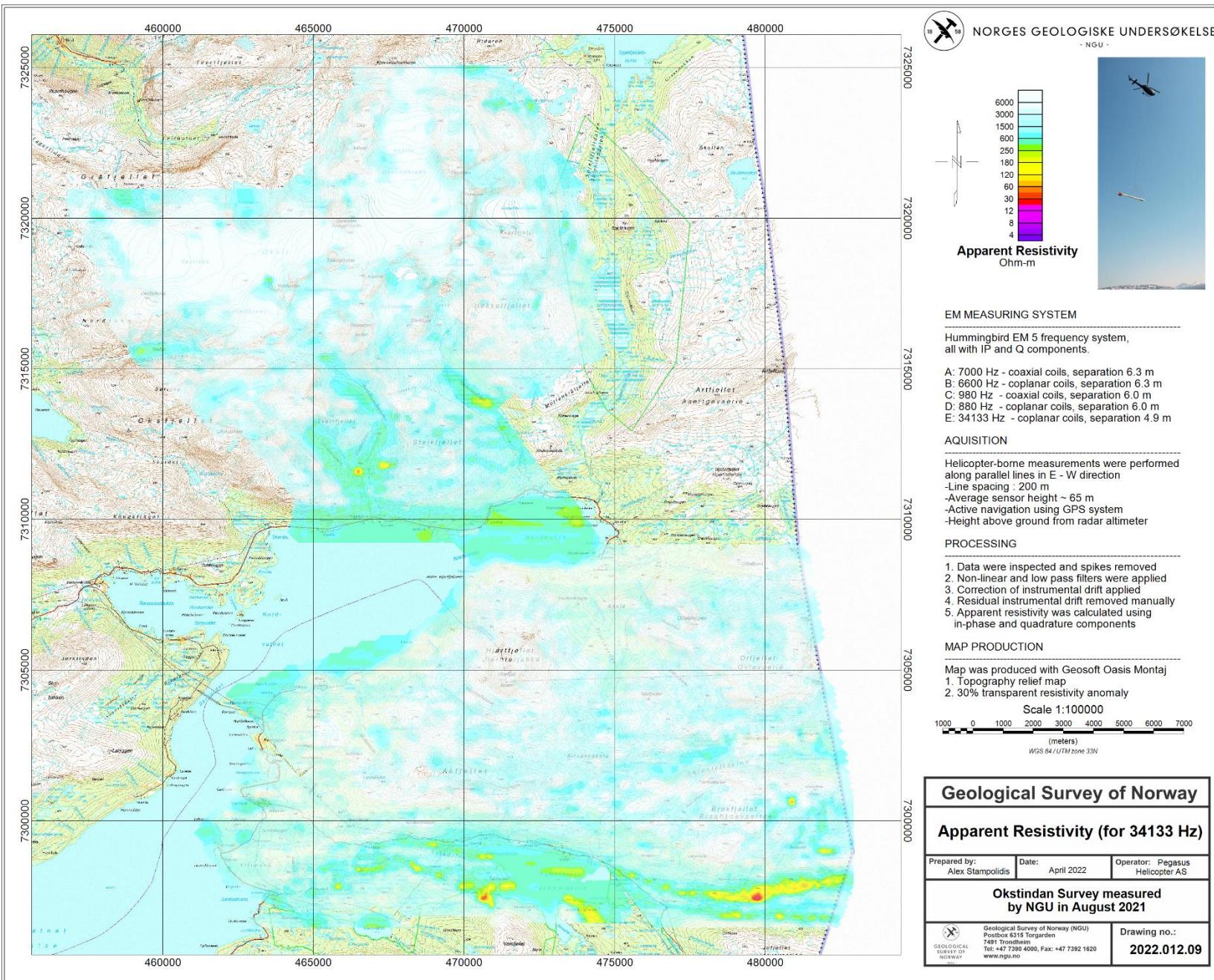


Figure 13: Apparent resistivity. Frequency 34133 Hz, Coplanar coils

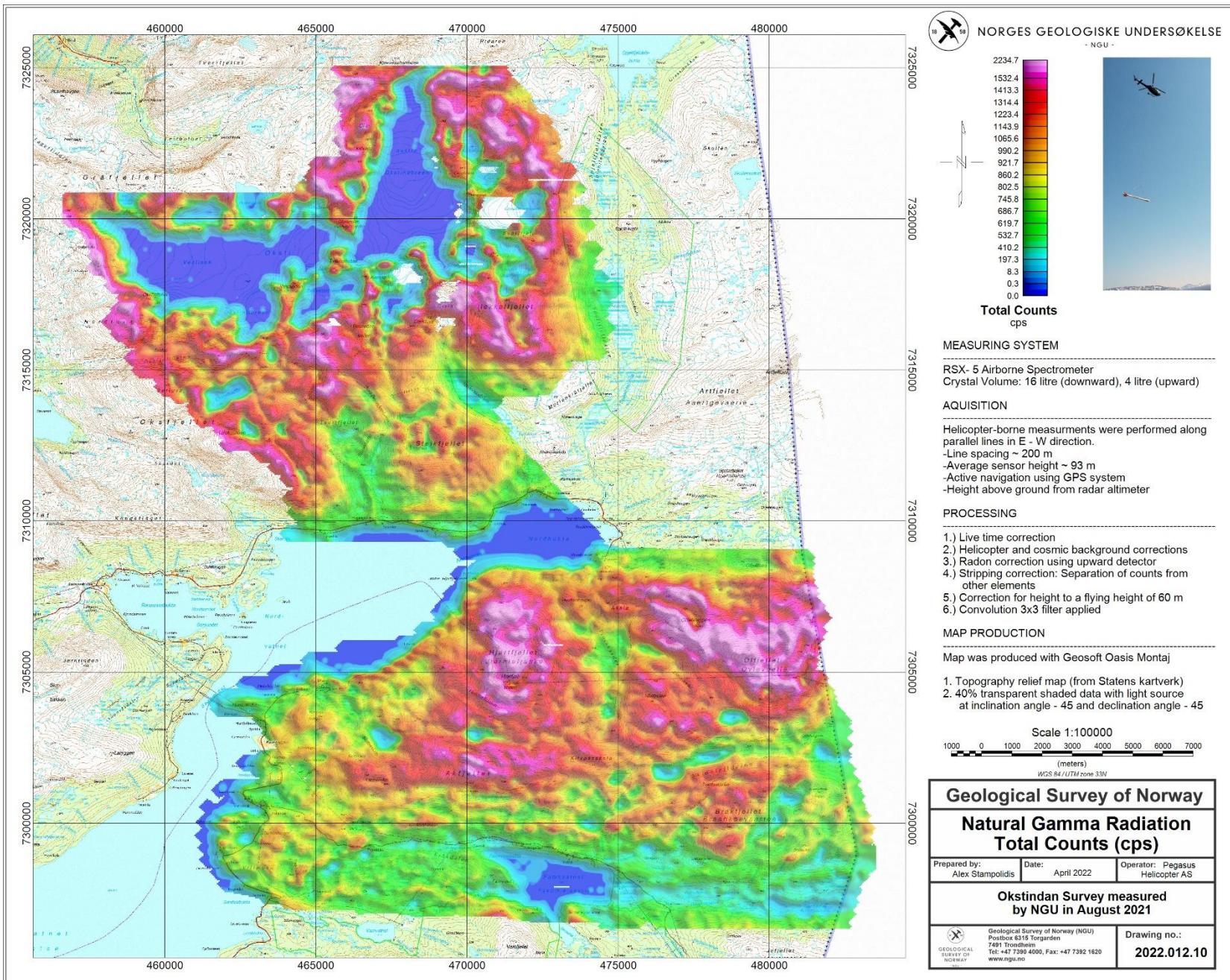


Figure 14: Radiometric Total Counts

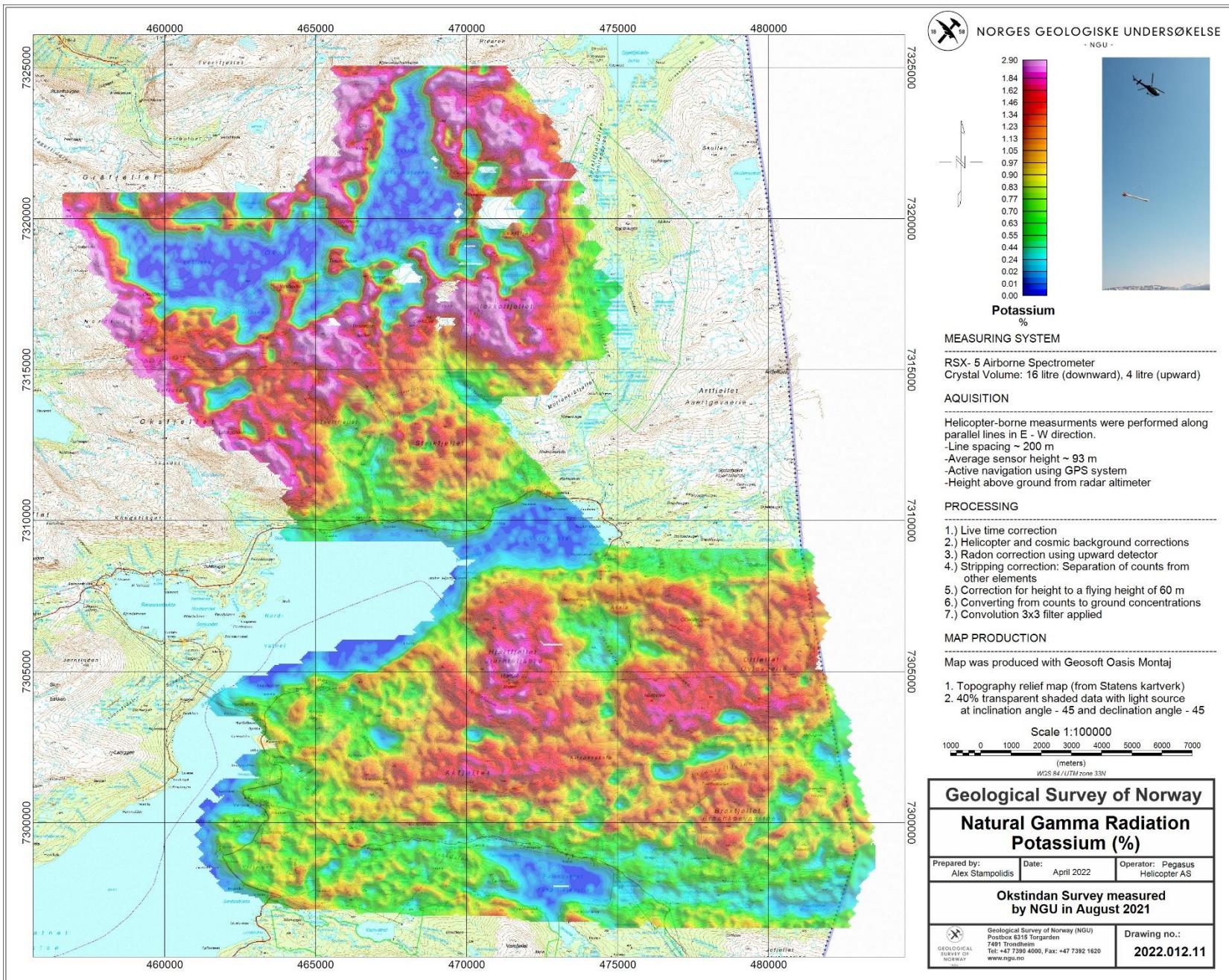


Figure 15: Potassium ground concentration

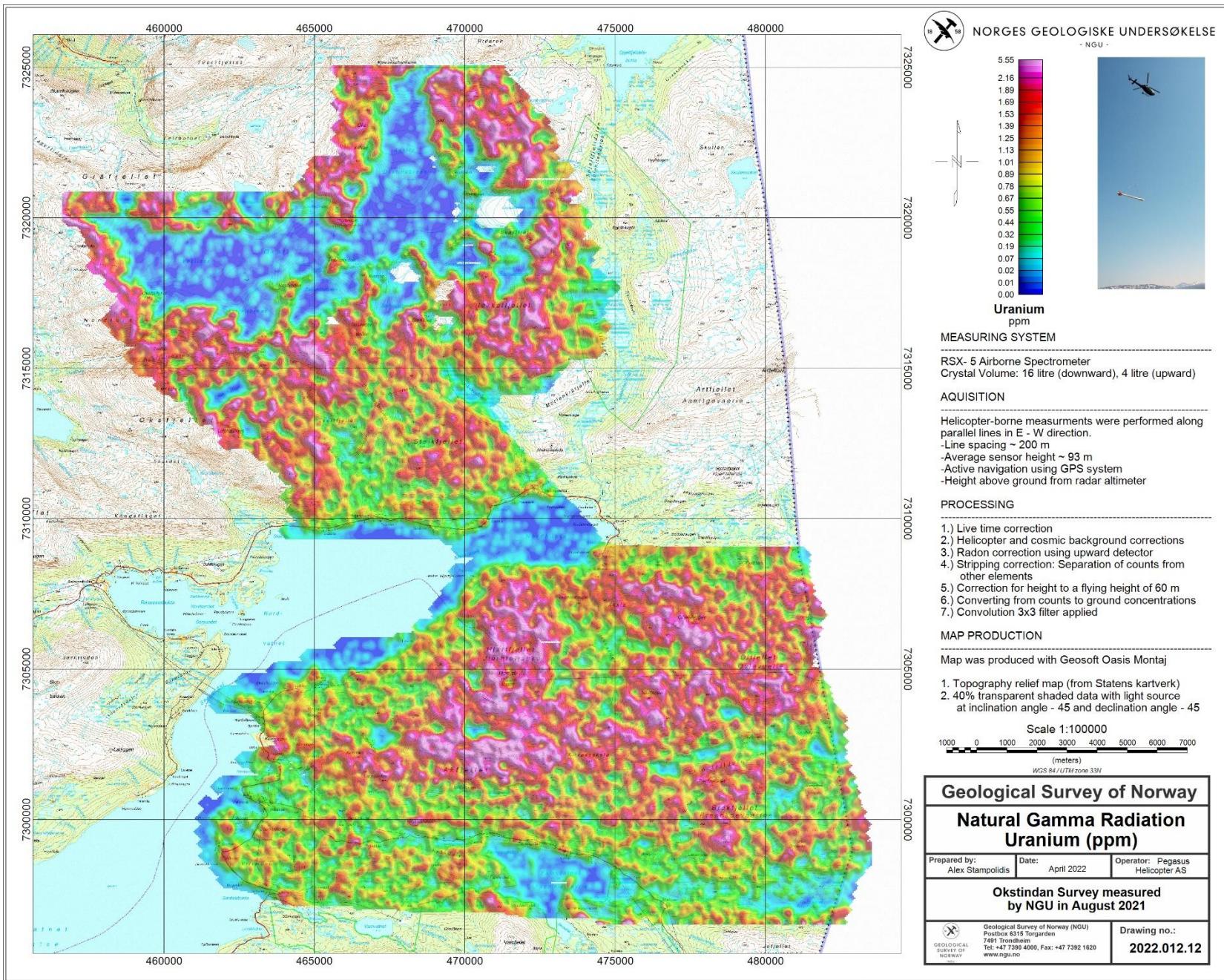


Figure 16: Uranium ground concentration

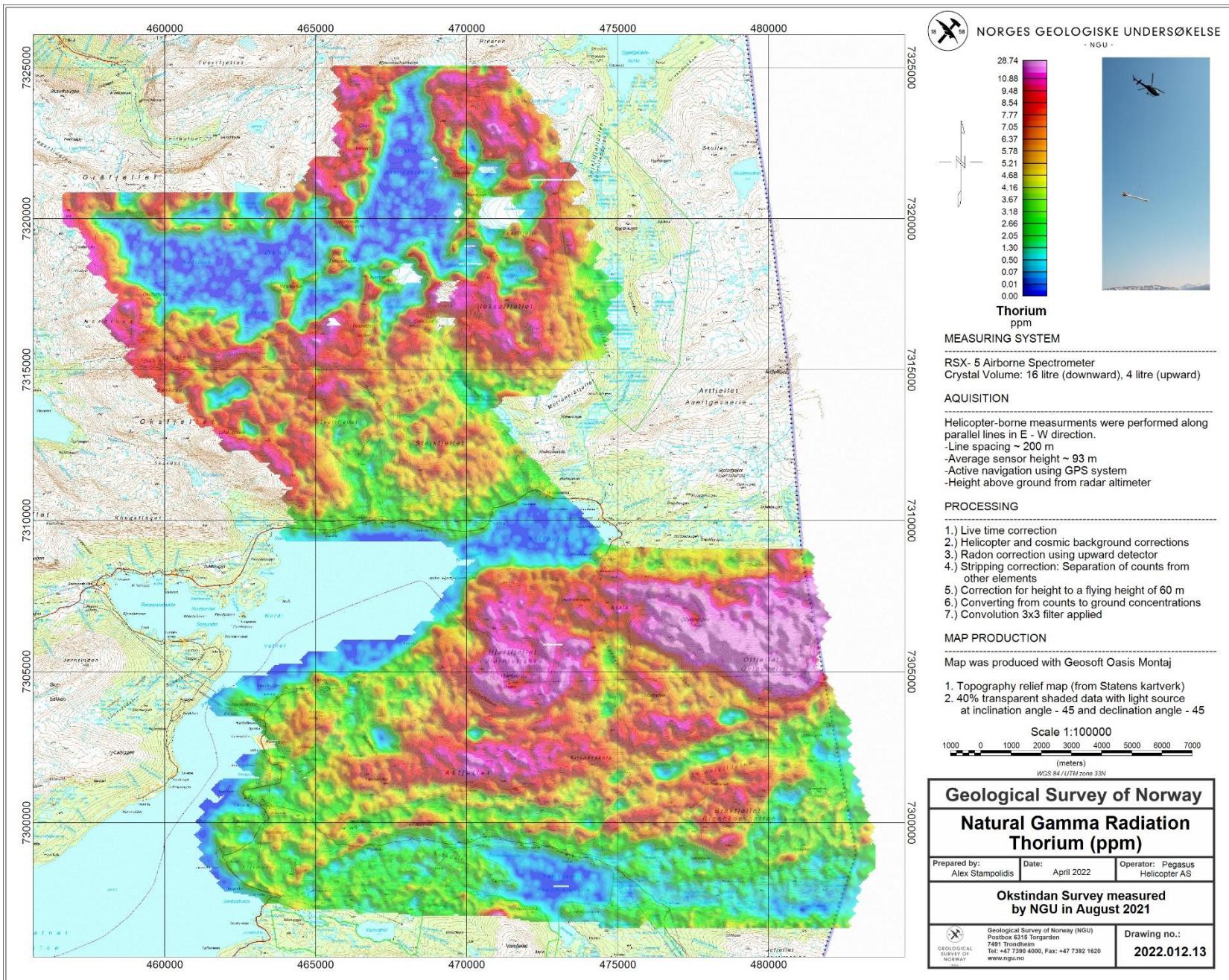


Figure 17: Thorium ground concentration

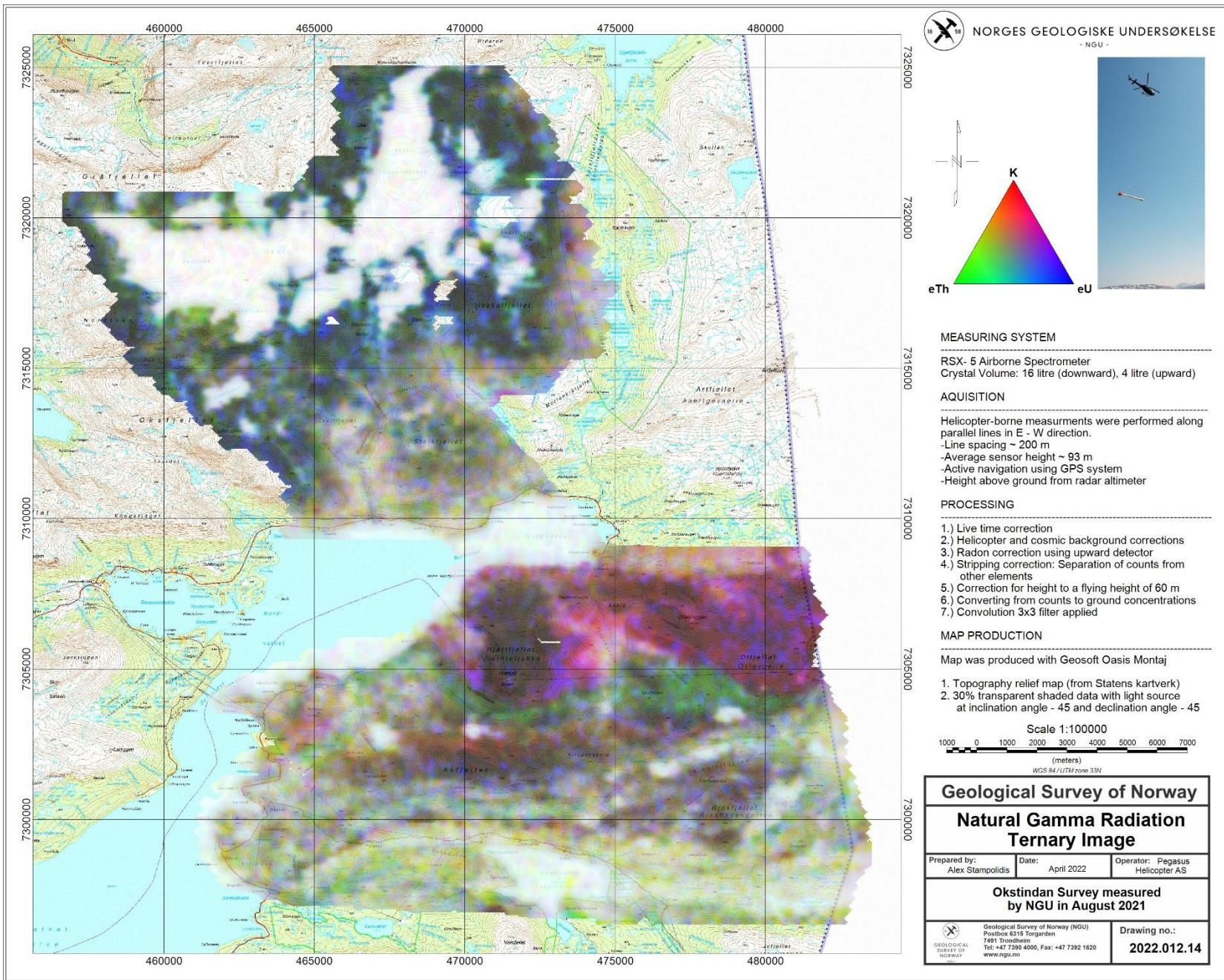


Figure 18: Radiometric Ternary Image



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